

cations and anions interrelationships during salinization and alkalization processes in some soils of a.r.e

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This study was carried out to investigate salinization and alkalization processes in connection with movement (downward or upward) of saline water through soil columns 30-cm in height and 6 cm in diameter. Downward movement was influenced by surface additions whereas upward movement was influenced by contacting bottoms of the soil columns with saline water in a form of NaCl solution. An empirical equation based on Gapon's equation for cation exchange equilibrium and similar to that of the U.S. Salinity Laboratory was established to deduce "ESP" as a diagnostic value of salt-affected soils by means of SAR of soil extract at 1:5 soil-water ratio. The reliabilities of applying the equation obtained in this work and three other equations developed by U. S. Salinity Laboratory (Richards, 1954), Balba (1962), and El-Arquan and Kaoud (1981) were examined. Other two equations were developed to deduce "ESP" through soluble ($\text{CO}_3 + \text{HCO}_3$) and pH of the 1:5 soil water extract. To investigate the possibility that soluble anions might in some way influence the cation exchange equilibrium between adsorbed and soluble cations, samples of three soils and two clay minerals (presaturated with Ca or Mg) were equilibrated with solutions of Na in the form of Cl or SO_4 . Also, samples presaturated with Ca were equilibrated with solutions of Mg in the form of Cl or SO_4 . The preference of the investigated samples for Na, Ca, and Mg was evaluated through calculating Vanselow's selectivity coefficient for Na-Ca, Na-Mg, and Mg-Ca systems. The results may be presented as follows:

1. With downward movement, soil salinity as well as soluble exchangeable Na were highest in the top layer of the soil column (0-10 cm) and lowest in the deepest one (20-30 cm). On the other hand, soluble and exchangeable Ca, Mg, and K increased with depth. Bicarbonates and chlorides underwent a pattern of distribution similar to those of soil salinity and soluble Na, whereas sulfates were comparatively lower in the top layer than in the subsurface one. Increasing concentration, as well as, the number of applications of the saline water increased soil salinity, soluble and exchangeable Na, soluble Cl and HCO_3 , but, in the same time, decreased soil content of soluble and exchangeable Ca, Mg, and K and did not affect soluble SO_4 .
2. With upward movement of saline water, salinity, as well as, soluble Na were highest in the top layer, lowest in the middle layer and intermediate in the bottom one. However, exchangeable Na and consequently exchangeable Na percentage were lowest in the surface and highest in the deepest layer. Soluble, as well as, exchangeable Ca, Mg, and K were highest in the surface and decreased gradually with depth. Chlorides accumulated in the top layers in amounts far greater than the layers below. Soluble bicarbonates underwent an opposite pattern, whereas soluble sulfates were distributed within the soil column in a pattern similar to those of soil salinity and soluble Na. Increasing concentration of the saline water and duration of its contact with the soil increased salinity of the soil, soluble and exchangeable Na, soluble Ca, Mg, K, and Cl. On the other hand, increasing input of salt decreased soil content of exchangeable Ca, Mg, and K, and did not affect soluble HCO_3 or SO_4 .

Deducing the ESP showed the following equations: (a) $\text{ESP} = 100 (0.0056 + 0.0311 \text{ SAR})$ (b) $\text{ESP} = 0.9299 + 9.5037 (\text{CO} + \text{HCO})$ (c) $\text{ESP} = -92.0418 + 12.1571 \text{ pH}$ Correlating values of ESP calculated from equation (a) as well as the other three similar equations developed by U.S. Salinity Laboratory (Richards, 1954), Balba (1962) and

El-Arquan and Kaoud (1981) with actually determined ESP supported the validity of using SAR for predicting ESP_c of the soil with a reasonable degree of success. Correlations between determined ESP and each of soluble ($CO_3 + HCO_3$) and pH were highly significant indicating the suitability of these parameters for predicting ESP. 4. There is specific anion effect on cation exchange equilibria. The adsorbed Na was greater when the associated anion was CO_3 rather than Cl or SO_4 . Adsorbed Na originated from the Na_2SO_4 solution was greater than that of the NaCl one. The variations became greater with increasing salt concentration. The selectivity coefficients calculated in terms of concentration " K_v " were higher with SO_4 as associated anion than with Cl. The differences became less when selectivity coefficients were calculated in terms of activity (K_{vt}) but were not eliminated. For all the samples, the Mg-saturated systems adsorbed more Na than the Ca-saturated ones. Also, selectivity coefficients of the former systems were always higher than the latter ones. At the values of SAR, the corresponding values of ESR in the case of SO_4 systems were higher than those in the case of Cl systems. When SAR_c (calculated on basis of activity) was used instead of " SAR " (calculated on basis of concentration) the differences were reduced but did not disappear. Adsorption patterns of Mg were similar to those of Na, however, at any equivalent fraction of the cation (Na or Mg) in the equilibrium solution the corresponding equivalent fraction of adsorbed cation in the Mg-Ca systems was higher than in Na-Ca or Na-Mg ones, consequently the preference of the investigated sample for Ca, Mg, and Na followed the order, $Ca > Mg > Na$.